

MACROFOSSILS AS INDICATORS OF PLIO–PLEISTOCENE CLIMATES IN TASMANIA AND ANTARCTICA

by R.S. Hill and G.J. Jordan

(with one table and five text-figures)

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ISSN 0080–4703. Department of Plant Science, University of Tasmania, GPO Box 252-80, Hobart, Tasmania, Australia 7001.

Plant macrofossils can be used as proxy palaeoclimatic indicators, since both the taxa present at a site and the form of leaves depend on the climate, and macrofossils are rarely transported far. In the simple case of the Sirius Group *Nothofagus* leaves and wood in Antarctica, the fossils provide data on climate that suggest temperatures at the time of deposition were in the order of >10°C warmer than at present. While the age of the fossils is still debated, the climatic signal is not. However, in more subtle cases, where the climatic differences are smaller and the fossil evidence is more complex, it is critical to interpret the evidence cautiously. Macrofossils from Pleistocene localities in Tasmania provide a clear signal on temperature and rainfall that is sometimes at odds with pollen-based reconstructions, but the former is inherently likely to be more accurate.

Key Words: plant macrofossils, Tasmania, Antarctica, climatic record, Neogene, Pleistocene.

INTRODUCTION

One approach to the assessment of palaeoclimates from plant fossils is to identify the fossils as accurately as possible, and then to use the distribution and ecology of the nearest living relatives (NLRs) to assess the climate. NLR estimates of Tasmanian palaeoclimates have been both qualitative (e.g. Macphail 1979 for microfossil data; Jordan *et al.* 1995 for the use of combined macro- and microfossil data) and semi-quantitative (e.g. Markgraf *et al.* 1986), particularly using palynological data. Macrofossils in Late Pleistocene deposits can usually be ascribed to species which still live in the region under consideration. The pollen and spores can usually be ascribed to extant genera, and sometimes to species. However, the specific identity of older fossils is less certain. Thus, most Early Tertiary species are extinct, some Early Pleistocene taxa are extinct (e.g. Jordan 1992, 1995a, b, Jordan & Hill 1994), but there are few records of extinct species in Late Pleistocene sediments (e.g. Jordan & Hill 1991). NLR methods may be less reliable if the species is extinct, since there is considerable variation in the ecological preferences of species within living genera. This problem is partly alleviated by combining the evidence from many fossil taxa within deposits, by using broad generic climatic preferences, and by identifying taxa with anomalous climatic preferences from other sites. The NLRs of macrofossils have some advantages and disadvantages over the NLRs of pollen and spores. Macrofossils can often be identified to higher resolution than palynomorphs, but there are often stronger taphonomic biases working on macrofossils than palynomorphs.

Fossil leaves and wood can contain climatic information other than that inferred from their taxonomic affinity, because they respond directly to climate. Both the gross morphology and the anatomy of leaves appear to respond to climate. This applies both within taxa (e.g. Jordan & Hill 1994), and also, by averaging over all the taxa or all the individual leaves, within sites (e.g. Christophel & Greenwood 1989, Wolfe 1993, Wing & Greenwood 1993). The models predicting climates averaging over all taxa or

all leaves are quite coarse in their resolution and of little use with Plio-Pleistocene floras (G.J. Jordan, pers. obs.). The use of the variation within species may provide some refinement of NLR climatic estimates (e.g. Jordan & Hill 1994). Similarly, wood anatomy (dendrochronological methods) can only provide coarse climate estimates when applied independent of the taxonomy, but can provide higher resolution when applied within species (Schweingruber 1988).

Well-preserved leaves and fruit of a wide range of extant and extinct species occur in a number of Pleistocene (ages follow Harland *et al.* 1990) sites in western Tasmania, and leaves and wood of a single species of *Nothofagus* occur in the probable Pliocene Sirius Group sediments in Antarctica. The present paper uses both NLR and morphological/anatomical data to assess Pleistocene climates from Tasmania and Pliocene climate from Antarctica.

SITES

Antarctica

The fossil material occurs in the Meyer Desert Formation of the Sirius Group at Oliver Bluffs, in the Dominion Range, Transantarctic Mountains, approximately 500 km from the south pole and 200 km inland from the junction of the Beardmore Glacier and the Ross Ice Shelf. The present elevation is approximately 1760 m above sea level (a.s.l.), but the mountains may have been uplifted to this altitude during the past two million years (Webb & Harwood 1993). This site is currently at 85°S and was at a similar latitude during the Pliocene. Plant macrofossils are preserved within several levels of the sediments, which consist of diamictites, sandstones and mudstones, interpreted to be the products of glacial, fluvial and lacustrine environments (Webb & Harwood 1993). This glacial material was carried by ice sheets from East Antarctica and deposited along the walls of a pre-Pleistocene valley system which bisected the Transantarctic Mountains.

The age of the sediments at Oliver Bluffs has been constrained by the presence of recycled marine diatoms in the glacial diamictites of the lower marine part of the formation. The youngest stratigraphic age of these diatoms suggests that the Sirius Group is younger than Late Pliocene in age (3.1–2.5 Ma) (Harwood 1983, 1986a, b, Webb & Harwood 1991). Radiogenic dating of volcanic ash in equivalent marine sediments in the CIROS-2 borehole in McMurdo Sound supports these dates (Barrett *et al.* 1992).

However, the dating of the sediments at Oliver Bluffs is still not clearly resolved. There have been claims that the diatoms may have been wind-blown in origin and trapped in fractures in the sediments, thus invalidating the diatom dates (e.g. Burckle & Pokras 1991). However, *Nothofagidites* pollen of the same species that is known from the Sirius Group has been reported in marine sediments from Cape Adare (DSDP Site 274, Ross Sea) of mid Pliocene age, supporting the view that *Nothofagus* trees were present on Antarctica at this time (Fleming & Barron, in press).

Tasmania

The Marionoak Formation at the Huskisson/Marionoak divide (41°41' S, 145°31' E, 430 m a.s.l.) contains glacial sediments overlain by fine-grained lacustrine sediments, including microfossils and some macrofossils. Palynostratigraphy implies that these sediments are likely to be Pleistocene and are no younger than the early part of the Middle Pleistocene (Macphail *et al.* 1993). However, these sediments are magnetically normal and underlie magnetically reversed sediments, implying that they are at least as old as the Jaramillo event in the Early Pleistocene; they are considered to be Early Pleistocene (Augustinus *et al.*, in press).

Clasts containing well-preserved macro- and microfossils occur in Middle–Late Pleistocene gravels at Regatta Point, western Tasmania (42° 10' S, 145° 20' E, 10 m a.s.l.) (Hill & Macphail 1985). Macphail *et al.* (1993) assigned an Early–?Middle Pleistocene age to a number of these clasts based on palynostratigraphy. One clast (referred to as RPU2) is palynostratigraphically at least as old as the Huskisson/Marionoak sediments and probably older. This clast is, therefore, of Early Pleistocene age. This RPU2 clast contains three thin, intensely fossiliferous layers. The composition of each layer differs little from the others. Other clasts (collectively referred to as RPU1 and including the clast described by Hill & Macphail 1985) are palynostratigraphically of same age as the Huskisson/Marionoak sediments (Macphail *et al.* 1993) and are likely to be of Early Pleistocene age. One clast (RPA) is palynostratigraphically younger than the Huskisson/Marionoak site and is considered to be of Early–Middle Pleistocene age (Jordan *et al.* 1995).

The Regency Formation in western Tasmania (42°10'S, 145°38'E, 180 m a.s.l.) consists of a 0.8 m thick layer of highly organic sediments, much of which is almost entirely composed of fossil leaves, fruit and twigs, and overlies an Early Pleistocene till. The formation was considered by Fitzsimons *et al.* (1990) to be of Early–Middle Pleistocene age, but is more likely to be of Middle Pleistocene age (E.A. Colhoun, pers. comm., Jordan & Hill 1994). The parts of the Langdon River sediments (41°58'S, 145°32'E, 520 m a.s.l.) discussed here are highly organic sediments made up of compressed beds of leaves and twigs (Colhoun *et al.*

1989). These represent some Middle Pleistocene interglacial, although a Last Interglacial age cannot be excluded at present (Colhoun *et al.* 1989).

METHODS

Climatic estimates were made from Tasmanian Pleistocene leaf macrofossil data, using the ranges in climatic parameters of the nearest living relatives of macrofossils. For Australian taxa, climatic parameters were estimated using ESOCLIM (H. A. Nix, J.P. Busby, M.F. Hutchinson & J.P. McMahon pers. comm.; Hutchinson 1991) analysis of a large data set of compiled point distributions (Floyd 1990; Environmental Resources Information Network; Department of Environment and Land Management, Tasmania; Tasmanian Herbarium; Australian National Herbarium; Australian National Botanic Gardens Herbarium). Distributions for New Zealand taxa were derived from Poole & Adams (1986), M.S. McGlone (pers. comm.) and G.J. Jordan (pers. obs.). Approximate climatic ranges for the New Zealand taxa were then derived from local climatic averages (New Zealand Meteorological Service 1981), allowing for an altitudinal lapse rate of 0.65°C/100 m. Unfortunately, more detailed climatic or distributional data was not available for New Zealand, but this does not effect the results of this work. The climatic ranges of the NLRs within sites were then compared graphically by simultaneously plotting ranges of climatic parameters of NLRs for all fossil taxa, and looking for overlaps. The region of overlap is assumed to be the best range of climatic estimates. Extreme outlier taxa were obvious and were excluded from the estimates. This method is similar to that of Kershaw & Nix (1988) and is reasonably robust, since random subsets of the data provide essentially the same estimates (i.e. supporting the method), and the method allows for taxa to be present near the ends of their ecological range. It also allows for the identification of outlying taxa, which may represent extinct species, extinct ecotypes or taxa whose modern ranges underestimate their potential climatic ranges.

Climatic estimates for the Sirius Group were prepared using qualitative methods, which will be explained in detail in the text below.

PALAEOCLIMATES

Sirius Group

The most controversial feature of the palaeoclimate of the Sirius Group is temperature, and the macrofossils present can be used to place clear constraints on this. Three aspects of temperature will be discussed: mean annual temperature, minimum winter temperatures, and mean summer maximum temperature. In order to demonstrate how these temperatures were estimated, some features of the macrofossils must be summarised.

(1) The leaves, wood and pollen in the Sirius Group sediments all belong to a single species of *Nothofagus* (Nothofagaceae), the southern beech. It is reasonable to presume that it was the same species, that is, that one biological species contributed all three organs to the sediments. The leaves have been formally described as *N. beardmorensis* (Hill *et al.*, in press), and that name will be used for the species forthwith.

(2) *Nothofagus beardmorensis* was a winter deciduous species (Hill & Truswell 1993, Hill *et al.*, in press). This is usually difficult to prove for a fossil, but in this case the evidence is very convincing.

(3) *Nothofagus beardmorensis* was a ground-hugging shrub, with no main axis, but many small, twisted branches with distinct, but very small growth rings (Francis & Hill in press). There is abundant evidence for abrasion on the wood, indicative of the effect of wind-borne ice and sand particles (Francis & Hill, in press).

(4) The macrofossils have not been transported more than a few metres from the growth position of the parent plant, and they have not been reworked. This is clear because of the thousands of delicately preserved leaves in a dense mat which could not have been reworked, and which certainly did not travel a long distance to be deposited in such dense numbers. It is also highly unlikely that leaves, wood and pollen, which have such different dispersal properties, could occur in the same sediments in such numbers if they had been transported long distances or, indeed, if they had been reworked. This is an important point because, although there is still doubt about the age of the fossils, there can be no doubt that the plants were growing on the site when the sediments were forming, and thus are a reliable guide to the prevailing climate, whenever it was.

From a combination of these features, estimates of temperature were made.

Minimum winter temperature

There are more than 30 living species of *Nothofagus*, but only seven are winter deciduous, six in South America and one in Tasmania. Some research has been done to estimate the frost tolerance of these species, and this has been summarised by Hill & Truswell (1993). Frost tolerance is likely to be a major factor limiting the distribution of a species as far as minimum winter temperatures are concerned. Direct frost damage occurs when ice crystals form within cells and lacerate the cell membranes. The damage this causes is irreversible, and the cell dies. Clearly, if this is widespread through the plant tissues, the plant will not survive. Frost damage does not necessarily occur at the same temperature in all organs and, for example, is unlikely to be relevant in the leaves of a winter deciduous species, since leaves are absent during the coldest part of the year. According to the data summarised by Hill & Truswell (1993), the lowest frost survival temperature yet recorded for *Nothofagus* is -22°C , and this occurred in a winter deciduous species (*N. antarctica*). This is likely to be a conservative estimate, since relatively few data have been collected for extant species, and *N. beardmorensis* is likely to have had a frost resistance at the low end of the extant range and possibly well beyond it. It should also be recognised that the temperature at which frost death occurs is not necessarily a good estimate of minimum winter temperature, since many plant species are known to be insulated against extreme cold in winter by a thick covering of snow. That is, if *N. beardmorensis* was covered in snow during winter, then temperatures could have fallen well below its frost death temperature without the plants being subjected to them. For these reasons, the estimate of -22°C for minimum winter temperature should be considered as a very conservative estimate.

Mean summer maximum temperature

A crude but very effective limit can be applied to summer temperatures — in order for a vascular plant to grow it must

have continued access to liquid water. If air temperatures rise above freezing, but the ground stays frozen, it is possible for a plant to transpire itself to death, as it loses what liquid water it has in its tissues but cannot replace it from the frozen ground. Clearly, at the time that *N. beardmorensis* was growing, summer temperatures were sufficient to thaw soil water to some reasonable depth to allow steady transpiration. This requires temperatures well above 0°C . Furthermore, in order to produce a crop of leaves and the reproductive structures required to produce the pollen present in the sediments, it is reasonable to assume that the daily maximum temperature must have been at least 5°C for several weeks. This is a conservative estimate based on woody plant growth requirements at high northern and southern latitudes today. It is probable that there were no particularly low minimum summer temperatures, since there is no evidence of frost damage on any of the leaf fossils.

Mean annual temperature

Francis & Hill (in press) argued that the wood macrofossils in the Sirius Group were very similar in form and function to the wood of the extant Arctic willow (*Salix arctica*), although there is no close taxonomic relationship between them. In fact, this model seems very precise, and since *S. arctica* grows further north than any other woody plant today, the lowest mean annual temperature at which it can grow provides a good physiological constraint on the peculiarities of wood structure observed in the fossil *Nothofagus*. Francis & Hill (in press) note that this mean annual temperature is approximately -15°C .

Comparison with the Sirius Group today

In order to understand the significance of these temperature estimates, it is useful to have some indication of comparable temperatures at the fossil locality today. No direct measurements exist, but approximations can be made by applying a lapse rate ($-0.65^{\circ}\text{C}/100\text{ m}$) to data from the nearest climate station. This has been done (D.M. Harwood, pers. comm.), but these estimates may contain considerable error. Nevertheless, the difference between these estimates and those from the fossils (table 1) are so great that any errors of calculation of extant temperatures must be small by

TABLE 1
Temperature data for Vanda Station* and estimates for the fossil location today†

	Vanda Station	Fossil location (extant)	Fossil location (extinct)
Mean summer maximum temperature	+2°C	-10°C	+5°C
Minimum winter temperature	-36°C	-48°C	-22°C
Mean annual temperature	-18°C	-28°C	-15°C

* $77^{\circ}32'S$ $161^{\circ}33'E$ near L. Vanda.

† Based on a lapse rate of $0.65^{\circ}\text{C}/100\text{ m}$ altitude (data provided by D. M. Harwood). Estimates for the fossil location at the time when *Nothofagus beardmorensis* was growing are discussed in the text.

comparison. The differences between mean summer maximum temperature and mean annual temperature are 15°C and 13°C respectively, with the site being warmer in the past, but the close correspondence between these two figures is probably fortuitous, given the inherent errors in the method. The difference between minimum winter temperature now and in the past at the fossil site is higher (26°C), but it was stressed earlier that the estimate for the fossils was likely to be conservative, thus inflating the difference.

Therefore, as a first approximation, a temperature difference in the order of 13–15°C occurred between the present day and the time at which the fossil plants were growing. Other factors may enter into this; for example, the possibility of post-depositional uplift of the fossil locality (Webb & Harwood 1993), which would artificially increase the temperature differential, although such uplift seems unlikely.

The other feature of interest in table 1 is that the difference in temperature between Vanda Station today (77°32'S 161°33'E, near Lake Vanda and about 800 km from the fossil site on about the same longitude) and the requirements for the fossil species is not very great at all, being, with the exception of the conservative minimum winter temperature, only 3°C. This is important, since it is usually considered that Antarctica is far too cold to support woody plant growth; that is not the case, and this may explain one of the dilemmas facing those arguing for the continual habitation of Antarctica by *Nothofagus* until the Pliocene. In fact, there is no reason for *Nothofagus*, in the time leading up to a Pliocene warming, not to have survived periods almost as cold as those that occur today, without recourse to unlikely long-distance dispersal of seed across prevailing winds.

Tasmania

Early Pleistocene

The macrofossil-bearing sediments at the Huskisson/Marionoak Divide, the overlying tills and the intervening slope deposits form a sequence representing the onset of full glacial conditions of the Bulgobac Glaciation. The Bulgobac Glaciation probably occurred at the same time as the Linda Glaciation in the King River catchment (Augustinus *et al.*, in press). This glacial probably produced the most extensive Pleistocene glaciation in Tasmania. The fossils in the Marionoak Formation, therefore, indicate conditions leading into a major Early Pleistocene glacial. Microfossils suggest that the vegetation was similar to subalpine woodland which now grows on the eastern Central Plateau of Tasmania. If correct, this indicates a rainfall in the order of 750–1000 mm per annum, which is 1000–1500 mm below current conditions at the site (Augustinus *et al.*, in press). This vegetation type now occurs in climates about 1–4°C below current conditions at the site (i.e. mean annual temperatures of about 6°–9°C). The Marionoak Formation contains leaves of *Orites acicularis* and *O. revoluta* (Jordan 1995a), two species of *Epacris/Archeria* and one of *Styphelieae* (Jordan & Hill, in press), and probable cyperaceous seeds. The two *Orites* species are subalpine–alpine species, although their leaf form suggests that they grew in sheltered areas. The two species of *Epacris/Archeria* are distinct from any species now in Tasmania, but they have extremely small leaves (about 1.5–2 mm long by about 0.4 mm wide), and the only Tasmanian species of this group with leaves as small as this are *E. navicularis*, *E. petrophila* and extreme high-altitude

forms of *E. serpyllifolia*. All of these are restricted to full alpine conditions. The climatic envelope implied by the macrofossils in the Marionoak Formation (fig. 1) suggests that the mean annual temperature was probably in the range of 3°–7°C, which is lower than the upper part of the range of the estimate from microfossils. The macrofossils provide little evidence about precipitation at this site.

The rich macrofloras in the Regatta Point sediments help elucidate several details of the climate at the time of deposition of this important fossil site, although they raise questions about some of the climatic indicator taxa. The flora has been divided into three broad groups, two of which (known as RPU2 and RPA) form discrete floras, but the third (RPU1) is an artificial composite of floras sharing a number of common features. The interpretation of the RPU2 flora is mainly based on a single layer (referred to as “17–18 cm” although it is only about 5–6 mm thick), since this undoubtedly represents a short depositional period (a few years or less) and the composition is similar to that of the other layers.

The RPU2 flora is mainly composed of elements which now occur at low altitudes in the region, but also contains a group of taxa which no longer occur in lowland western Tasmania, including apparent warm climate indicators such as macrofossils of *Quintinia* and *Dacrycarpus* (Hill & Macphail 1985, Jordan & Hill 1991, 1994, Jordan 1992, 1995a, b, Jordan & Dalton, 1995). The temperature range envelopes of the macrofossils do not all overlap (fig. 2). However, all but four NLRs occur in sites with mean annual temperatures between 11° and 12°C, although this is at the extreme end of the range of several taxa. Any temperature estimate outside this range involves more discrepancies. The current mean annual temperature at Regatta Point is about 12°C. The four taxa which do not occur at mean annual temperatures of 11°–12°C are *Microcachrys tetragona*, *Richea milliganiilacerosa*, *Rubus nebulosus* and *Callitris macleayanus*. Both the *Rubus* and *Callitris* fossils are globally extinct species, and the discrepancies in temperatures are small. It is likely that these species tolerated cooler climates than their NLRs. The fossil *M. tetragona* has foliage indistinguishable from the extant species, but the distinctive pollen type is common in Tertiary deposits containing a wide range of floras, including many from

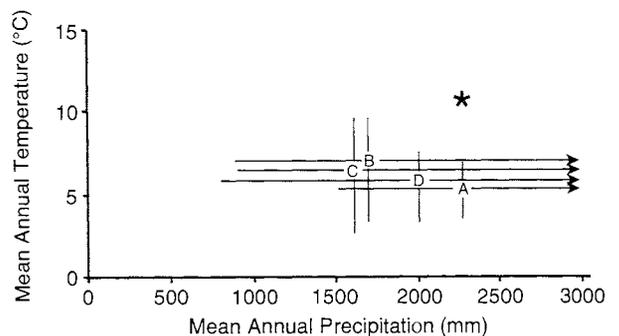


FIG. 1 — Current ranges in mean annual temperature and mean annual precipitation for the nearest living relatives (NLRs) of macrofossil seed plant taxa (A–D) found in the Early Pleistocene sediments at the Huskisson/Marionoak divide site. A = *Microcachrys tetragona*, B = *Orites revoluta*, C = *O. acicularis*, D = *Epacris petrophila/navicularis*. The asterisk indicates the current climate at the site.

undoubtedly much warmer climates than the modern range of this species. It is unlikely that the macrofossils of *M. tetragona* were derived from uplands areas (Jordan & Hill 1994; Jordan 1995b). The identification of the *Richea*-type fossil is relatively broad, and this may also represent an extinct species. A possible scenario to explain the apparent anomalies is that the Regatta Point sediments were deposited early in an early interglacial, with the warm climate indicators occurring in more fertile wet forest and cool climate taxa persisting in oligotrophic areas. It is common now for subalpine type vegetation to occur in oligotrophic sites at moderate altitude alongside more typical lowland vegetation on less oligotrophic areas, particularly in areas of abundant rainfall.

There is little local topography in the Regatta Point area, and the fossiliferous clasts are highly unlikely to have been transported far, so their original site of deposition was not significantly higher than their present location. There are no micro- or macrofossils in the flora which indicate an estuarine environment. The sediments are at low altitude (10 m a.s.l.) beside an estuary, so sea level at the time of deposition was not more than about 10 m above present sea level. This also argues against the climate being significantly warmer than it is today. Thus, it is highly likely that the mean annual temperature at Regatta Point at the time of deposition of the RPU2 sediments was similar to the modern mean annual temperature and precipitation.

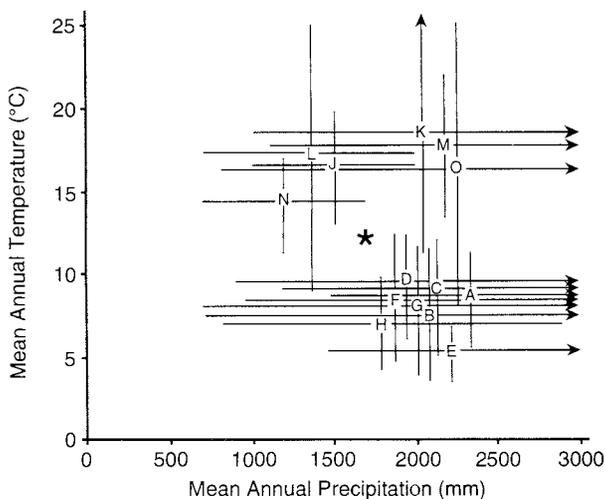


FIG. 2 — Current ranges in mean annual temperature and mean annual precipitation for the nearest living relatives (NLRs) of macrofossil seed plant taxa found in the Early Pleistocene RPU2 sediments at Regatta Point. A = *Agastachys odorata*, B = *Athrotaxis* spp., C = *Eucryphia lucida*/milliganii, *Trochocarpa gunnii*/*cunninghamii*, *Orites diversifolia*, D = *Lagarostrobos franklinii*, E = *Microcachrys tetragona*, F = *Nothofagus cunninghamii*, G = *Telopea truncata*, H = *Richea milliganii*/*acerosa*, J = *Rubus nebulosus*, K = *Lauraceae* (except *Cassytha*), L = *Banksia spinulosa*, M = *Callitris macleayanus*, N = *Oxylobium ilicifolium*, O = *Dacrycarpus dacrydioides* and *Quintinia acutifolia*. The asterisk indicates the current climate at the site. Some taxa, including *Acacia melanoxylon*, *Styphelioideae*, *Dicksonia antarctica*, *Epacris* spp., *Pomaderris apetala* and *Allocasuarina*, have been excluded from this plot because their climatic ranges are very broad and overlap those of several other taxa.

There is no evidence that the mean annual precipitation is any different from now, because all the precipitation of the NLRs overlap at approximately the current precipitation at Regatta Point.

The RPU1 sediments contain an assemblage of taxa which is broadly similar to that found in the RPU2 sediments, except that there are more taxa with cool climate NLRs and fewer with warm climate NLRs (Macphail *et al.* 1993). Thus, the fossils in these sediments are consistent with cooler climates than the RPU2 sediments, but the evidence is not strong. Also, the RPU1 sediments are a more or less artificial assemblage of clasts, and hence climatic inferences are weak.

The RPA sediments are about 0.05 m thick but are likely to represent only a short period, since the matrix is coarse sand, and rapid burial under these conditions is almost certainly necessary to allow the establishment of anaerobic conditions. The macro- and microfossils in the RPA sediments have been used to suggest that the vegetation was similar to subalpine vegetation (Jordan *et al.* 1995). Most of the macrofossil taxa can be confidently assigned to living species. All the taxa in the sediments are consistent with Tasmanian subalpine vegetation except for spores of *Cyathea*, and pollen of *Symplocos* and *Casuarinaceae* (fig. 3). These taxa are considered to be outliers, whose present distributions do not indicate their past preferences. *Cyathea* is restricted to small lowland pockets and *Casuarinaceae* does not occur at high altitudes, but both *Casuarinaceae* and *Cyathea* spores are widespread and common in western Tasmanian sediments older than the last glacial maximum, even among clearly cold climate assemblages. In addition, the *Allocasuarina* fossils in the RPU1 and RPU2 sediments are of an extinct species. The Huskisson/Marion Oak Divide flora also contains *Symplocos* pollen, which suggests that this pollen type in the RPU1, RPU2 and RPA sediments was produced by a species which grew outside the modern climate range of the genus. The climatic envelopes (fig. 1) indicate that the temperature was most likely to be within the range of 5°–8°C, and thus 4°–7°C cooler than at present. The climatic envelopes suggest that the precipitation was probably more than about 1200 mm per annum (fig. 1) and, thus, little, if at all, drier than at present.

Middle Pleistocene

There are significant changes in the composition of both macro- and microfossil assemblages through the Regency Formation (Fitzsimons *et al.* 1990). This has been interpreted as most likely showing a change in vegetation from cooler climates into full interglacial conditions, with probably a similar climate to that at present, although these changes may be the result of successional change (Fitzsimons *et al.* 1990). The climatic ranges of the NLRs of macrofossil taxa occurring in the lower part of the Formation (Zone RY1 of Fitzsimons *et al.* 1990) all overlap in the range of mean annual temperature 8.1–11.4°C (fig. 4). In the upper section (Zone RY2) the overlap is in the range of 6.1–11.6°C (fig. 5). Therefore, the climatic ranges of the NLRs of all the macrofossils throughout the sediments are consistent with modern temperatures. The climatic ranges of the NLRs of all the microfossil taxa in the sediments (Fitzsimons *et al.* 1990), except *Nothofagus gunnii* and *Microcachrys tetragona*, also overlap with the modern climate, and both *N. gunnii* and *M. tetragona* occur as rare grains both above and below the supposed climate change in the Regency sediments. Thus, the palaeobotanical evidence for climatic change in

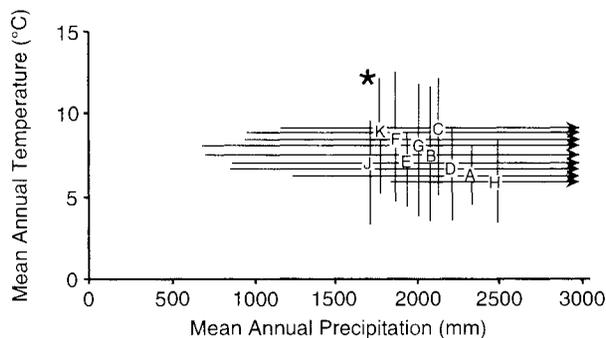


FIG. 3 — Current ranges in mean annual temperature and mean annual precipitation for the nearest living relatives (NLRs) of macrofossil seed plant taxa (A–J) found in the Early–Middle Pleistocene RPA sediments at Regatta Point. A = *Athrotaxis cupressoides*, B = *Athrotaxis selaginoides*, C = *Eucryphia lucida/milliganii*, D = *Diselma archerii*, E = *Nothofagus gunnii*, F = *Nothofagus cunninghamii*, G = *Telopea truncata*, H = *Orites milliganii*, J = *Orites revoluta*, K = *Phyllocladus aspleniifolius*. The asterisk indicates the current climate at the site.

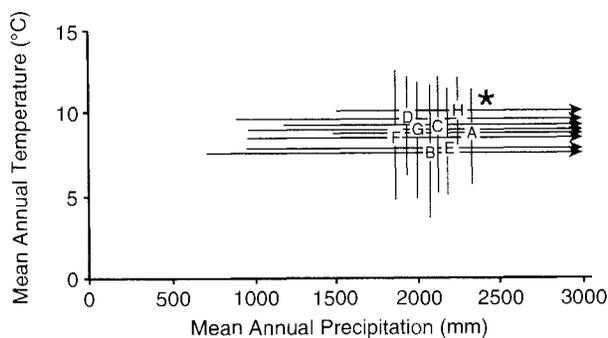


FIG. 4 — Current ranges in mean annual temperature and mean annual precipitation for the nearest living relatives (NLRs) of macrofossil seed plant taxa found in the lower section (Zone RY1 of Fitzsimons *et al.* 1990) of the Middle Pleistocene Regency Formation. A = *Agastachys odorata*, B = *Athrotaxis selaginoides*, C = *Trochocarpa gunnii/cunninghamii*, D = *Lagarostrobos franklinii*, E = *Eucryphia lucida*, F = *Nothofagus cunninghamii*, G = *Phyllocladus aspleniifolius*, H = *Allocasuarina zephyrea*. The asterisk indicates the current climate at the site.

this site is weak and is based on changes in the abundance of taxa. The change in composition of the flora could easily be explained by vegetation succession, or possibly by changes in the depositional environment.

Discussion of Tasmanian Pleistocene data

There is no convincing macrofossil or microfossil evidence of climates warmer than now in the Pleistocene in Tasmania. The best estimates of temperature for all sites are the same as, or lower than present climates. The fossils at the Early Pleistocene RPU2 sediments and the Middle Pleistocene Regency Formation are consistent with climates as warm as those which now prevail. There are other Middle and Late Pleistocene sites containing macrofossils, including the

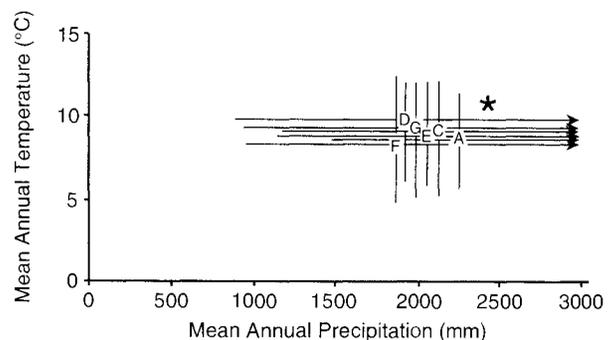


FIG. 5 — Current ranges in mean annual temperature and mean annual precipitation for the nearest living relatives (NLRs) of macrofossil seed plant taxa found in the upper section (Zone RY2 of Fitzsimons *et al.* 1990) of the Middle Pleistocene Regency Formation. A = *Agastachys odorata*, C = *Trochocarpa gunnii/cunninghamii*, D = *Lagarostrobos franklinii*, E = *Eucryphia lucida*, F = *Nothofagus cunninghamii*, G = *Phyllocladus aspleniifolius*. The asterisk indicates the current climate at the site.

Langdon River site (Colhoun *et al.* 1989) and Melaleuca Inlet (Jordan *et al.* 1991), which, based on similar arguments, are consistent with temperatures similar to modern temperatures. There are constraints on the upper estimates of climates for lowland sites younger than the Early Pleistocene, since virtually all the taxa involved are still extant in Tasmania and, hence, their climatic ranges include Tasmanian climates. However, it is also clear that many taxa previously used as warm climate indicators probably occurred in cooler climates than their NLRs, and their presence cannot be used as good evidence of warm climates. There is convincing evidence of colder climates than at present at times in the Early–Middle Pleistocene (1–4°C colder at the Huskisson/Marion Oak Divide site and 4–7°C colder in the RPA sediments).

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